The 2021 update to New Zealand's **National Tsunami Hazard Model**

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ABSTRACT

The National Tsunami Hazard Model (NTHM) provides estimates of tsunami heights at the shoreline for return periods of up to 2500 years. It is used as a basis for selecting scenarios used for tsunami inundation modelling, which underpins tsunami evacuation zone design, land-use planning and risk assessments. The original NTHM was completed in 2013, though it built upon a 2005 study focused only on the main coastal cities. Since 2013, there have been improvements in the understanding of tsunami sources and in the techniques of tsunami modelling, and this report describes an incremental update incorporating those improvements. This new revision we refer to as the 2021 NTHM.

One area of improvement is in the definition of earthquake recurrences, and the update uses parameters from the Global Earthquake Model (GEM), which was still in development at the time of the 2013 NTHM.

The 2021 NTHM makes use of a broad catalogue of simulated tsunami events caused by subduction earthquakes from around the Pacific. This catalogue was developed in 2019, initially for the purpose of informing real-time tsunami forecasts. By using this large catalogue of scenarios, and through more sophisticated use of scaling relationships, it is expected that the 2021 NTHM should produce more accurate tsunami hazard estimates.

Modelling of tsunami caused by crustal faults close to New Zealand has also been significantly improved. The 2013 NTHM used simple empirical relationships to estimate tsunami heights, and here those have been replaced by using scaled hydrodynamic models, which should produce significantly more accurate results.

Overall, we find that the 2021 NTHM tends to produce similar or slightly lower estimates of tsunami hazard at the shoreline compared to the 2013 NTHM at most locations and return times. The main exception to this is the east coast of the North Island, for which some areas are estimated to have higher hazard at longer return times. We anticipate that the introduction of improved tsunami height estimates will lead to better choices of scenarios for inundation modelling studies.

KEYWORDS

Tsunami, Hazard Model, Earthquakes, New Zealand

1.0 INTRODUCTION

The National Tsunami Hazard Model (NTHM) is a key science product for reducing the impact of tsunami in New Zealand by providing estimates of tsunami heights at the shoreline for return periods of up to 2500 years. It underpins a variety of other products, such as the evacuation zones used when a tsunami threatens New Zealand, probabilistic tsunami inundation assessments used in land-use planning and probabilistic tsunami risk assessments. The original NTHM was completed in 2013 (Power 2013). Since then, knowledge of tsunami sources and how to model resulting tsunami has improved. We have recently completed an incremental update to the NTHM, and this report outlines the changes that have been made. For the purposes of this report, this new revision will be called NTHM 2021, while the original model will be called NTHM 2013.

Some of the significant changes include:

- Updates to the subduction source earthquake recurrence estimates using the parameters from latest Global Earthquake Model (GEM) global subduction zone model. This is described further in Section [2.1.](#page-7-1)
- Updates to the scenario database used to estimate the shoreline wave heights of regional and distant source tsunami to make use of the most recent modelling used for producing tsunami threat maps following a large earthquake. Further details are in Section [2.2.](#page-9-0)
- Revision of the method used to estimate wave heights from local crustal tsunami sources, replacing a very simple empirical modelling approach with scaled hydrodynamical modelling. Further details are in Section 3.

The effects of the changes on estimates of shoreline tsunami hazard around New Zealand are shown at a national scale in Section [3.1,](#page-20-1) and tsunami hazard curves and deaggregation plots for individual cities are shown in Section 3.2.

2.0 METHOD

The following sections describe the process used to estimate the tsunami hazard around the New Zealand coast. As this report describes an update to the 2013 NTHM (Power 2013), the focus here is on describing the changes made to the 2013 model.

The hazard analysis results in estimates of tsunami hazard along the coast, which, for this purpose, has been divided into a set of coastal zones [\(Figure 2.1\)](#page-7-2).

Figure 2.1 Illustration of the 268 coastal zones that are used for defining the tsunami hazard around the coast. Each zone is approximately 20 km in length.

2.1 Revised Subduction Zones Parameters

Since the publication of the 2013 NTHM, further work has been done to better understand and characterise the likelihood of large earthquakes on the world's subduction zones. To update the new NTHM with this new information, we revised the source zones using the GEM (Global Earthquake Model) Faulted Earth Subduction Zone Characterisation Project's 2015 model (Berryman et al. 2015). The aim of that project was to develop a globally consistent characterisation of the world's subduction zone plate boundary faults. The project produced a globally consistent set of earthquake characterisation parameters for all subduction zones in the world. For the NTHM 2021, we updated the NTHM 2013 table of subduction zone parameters with the new information from Berryman et al. (2015), while retaining the subduction zone segmentation assumptions used in NTHM 20[1](#page-7-3)3¹, and then used these to calculate the revised shoreline tsunami hazard.

¹ Notably, NTHM 2021 continues to assume that the Hikurangi, Kermadec and Tonga subduction zones behave independently.

Most of the subduction zone parameters in NTHM 2021 are quite similar to those used in NTHM 2013. The most significant changes are typically in the b-value assumed for each subduction zone. The b-value is the slope of the magnitude versus frequency distribution for earthquakes following the Gutenberg-Richter equation. Higher b-values correspond to a smaller proportion of high-magnitude earthquakes compared to lower-magnitude ones. In both NTHMs, subduction zone parameters, such as b-value, are randomly selected between a minimum and maximum value using a uniform probability distribution. A particular set of sampled values are then used to create a synthetic catalogue of earthquakes. This process is then repeated using a new set of samples, generating numerous earthquake catalogues and thus estimates of the tsunami hazard. The spread of different hazard curves then reflects the uncertainties in the input parameters. This is explained in greater detail in Section 6 of the NTHM 2013 report (Power 2013). In NTHM 2013, the same range of b-values between 0.5 and 1.0 was used for all subduction zones. [Table 2.1](#page-8-0) shows the b-values used in NTHM 2021. Most zones now have their own minimum b-value, normally much higher than 0.5. The maximum b-value for most zones is 1.2, up from 1.0 in NTHM 2013.

The maximum magnitude range for each zone is similar in NTHM 2013 and 2021. One exception is the Alaska/Aleutian zone, where the range changes from M_w 9.5 to 9.7 in NTHM 2013 to 9.2 to 9.6 in NTHM 2021 (consistent with Berryman et al. 2015). For all other subduction zones, the maximum magnitude differs by less than 0.1 magnitude unit. The full list of maximum magnitude ranges for each zone is in Berryman et al. (2015). Coupling coefficients, plate velocity rates and directions and subduction zone lengths and widths were also similar or unchanged in both models. The ones used in this revision are those presented in Berryman et al. (2015) and so are not repeated here.

Subduction Zone Name	NTHM 2021 b-Value Minimum	NTHM 2021 b-Value Maximum	
Alaska/Aleutians	0.67	1.2	
Cascadia	0.7	1.2	
Japan/Kurile	0.61	1.2	
Kanto	0.7	1.2	
Nankai/Ryukyu	0.61	1.2	
Japan/Kurile	0.63	1.2	
Nankai/Ryukyu	0.61	1.2	
Izu-Bonin	0.7	1.2	
Marianas	0.68	1.47	
North Yap	0.7	1.2	
Palau-South Yap	0.7	1.2	
Hikurangi	0.7	1.2	
Kermadec	0.7	1.21	
Tonga	0.7	1.21	
Puysegur	0.7	1.2	
Hjort	0.7	1.2	

Table 2.1 Table comparing the minimum and maximum b-values used in NTHM 2021. Details of the location of the subduction zones can be found in Power (2013). In NTHM 2013, all zones used the same minimum and maximum b-value (0.5 and 1.0, respectively).

2.2 Regional and Distant Source Tsunami Heights

The 2013 NTHM used a collection of pre-calculated models of tsunami from a particular regional or distant source to estimate the coefficients in a semi-empirical scaling relationship based on Abe's (1979, 1995) formula:

$$
Ht_{ij} = 10^{M_w - B_{ij}}
$$
 Equation 2.1

where Ht_{ij} is the tsunami height at the coastal zone *j* due to an earthquake of moment magnitude *MW* in source region *i* (for distant and regional sources, the 'source regions' are the subduction zones) and *Bij* is a coefficient specific to tsunami travelling from source region *i* to coastal zone *j*. The *Bij* coefficients were derived from a catalogue of tsunami scenario models. The catalogue used in 2013 was the 2012 Tsunami Scenario Database (Power, unpublished; an earlier version of the database is described in Power and Gale 2011). This database was initially created to provide GNS Science with a rapid estimate of the tsunami threat level to its coastal warning zones following an earthquake on one of the major subduction zones around the Pacific.

For the 2021 NTHM, a revised and much larger set of scenarios has been used which were developed for an updated Tsunami Scenario Database developed in 2019 (Gusman et al. 2019, 2020). One of the main improvements was a significant increase in the number of scenarios in the new database. Scenarios now exist down to small magnitudes for all of the subduction zones near New Zealand. The number of scenarios increased to 998 (up from 336 scenarios in the [2](#page-10-2)012 database).² For more details about the update, see Gusman et al. (2019, 2020). Analysis of these scenarios indicated that the B_{ii} coefficients were not entirely independent of magnitude. For the current study, the average B_{ij} and their standard deviations were calculated at the specific magnitudes of the scenarios by rearranging [Equation 2.1](#page-9-1) as:

$$
B_{ij}(M_w) = M_w - log_{10}Ht_{ij}
$$

More specifically, $B_{ij}(M_w)$ was individually calculated for each scenario that occurs in source region *i*, and these were evaluated to produce a mean and standard deviation of $B_{ij}(M_w)$ for the source region *i* as a whole. For each subduction zone, the set of scenarios that arise from that particular source region was found by searching for those scenarios in the database that had rupture centroids within a polygon defining the boundaries of the subduction zone. This is a more flexible approach than that used previously and should make it easier to add more scenarios in future. The Ht_{ij} values were also determined for various percentiles of the tsunami height distribution along each section of coast using the new database, and then these values were used to calculate B_{ii} values for particular percentiles. This gives us the potential to create national-scale tsunami hazard maps for a range of different return periods and percentiles. The results presented here are based on the 99th percentile of the tsunami height distribution within each section of coast.

When used to estimate tsunami heights in the hazard model, interpolation was used to find a $B_{ii}(M_w)$ for the specific magnitude of each event in the synthetic catalogues. To allow for the effects of different locations within the source regions, the standard deviation $\sigma B_{ij}(M_w)$ is also estimated. The tsunami height is then calculated using [Equation 2.3,](#page-10-1) where $\overline{\sigma B_{11}(M_{yy})}$ is sampled from a normal distribution with mean zero and standard deviation $\sigma B_{ii}(M_w)$ and the normal distribution is truncated at +/- 2σ.

$$
Ht_{ij} = 10^{M_W - (B_{ij}(M_W) + \overline{\sigma_{B_{ij}}(M_W)})}
$$
 Equation 2.3

2.3 Local Subduction Zone Tsunami Height Estimation

For the local subduction zones, that is, those within 1 hour travel time to the New Zealand coast (i.e. the Hikurangi, Puysegur and Kermadec subduction zones), a similar method was developed to that used for the distant and regional sources. However, for these nearby sources, there is greater sensitivity of tsunami heights to the effects of earthquake location, and, rather than treating this sensitivity as a random input (via the standard deviation in B_{ij}), a different approach was used to find a $B_{ij}(M_w)$ that assumed a random position for the earthquake within the subduction zone.

² There were 723 scenarios in the first round of updates (Gusman et al. 2019) and a further 275 in the second round of updates (Gusman et al. 2020). Not all of these database scenarios were used in the NTHM analysis; in particular, a revised set were used for local subduction zones, see Section 2.3.

Figure 2.2 Illustration of how scenarios are used to obtain *B_{ij}* values for the Hikurangi and Puysegur subduction zones (the method for Kermadec Trench is similar but uses different magnitudes and more locations). A random position 1–12 is selected (yellow bar), then B_{ij} values from scenarios whose magnitudes bracket the event magnitude are selected (grey boxes represent scenario rupture areas) and a magnitude dependent *Bij(Mw)* is calculated from them by interpolation.

The scenarios used for this modelling were similar to those in the 2019 scenario database (Gusman et al. 2019, 2020), except that the scenario magnitudes and locations were chosen such that a whole number of rupture areas would fit within the subduction zones without gaps or overlap (as in [Figure 2.2\)](#page-11-1), giving equal coverage to all parts.

2.4 Local Crustal Faults Tsunami Height Estimation

NTHM 2013 used a set of empirical equations based on Abe's (1979, 1995) formula to calculate the amplitude of the tsunami at the coast (*Ht*):

$$
Ht = 10^{M_w - \log R + 5.55 + C}
$$
 Equation 2.4

In this equation, *R* is the distance between the earthquake's epicentre and the coast (in kilometres) and *C* is a constant between 0.0 and 0.2, depending on the location. For NTHM 2013, *C* was assumed to be 0.1 for all coastal zones around New Zealand. Since the above formula becomes large at small *R*, a different formula was used when *R* is small:

$$
Ht = 10^{0.5M_w - 3.30 + C}
$$
 Equation 2.5

Here, *R* is assumed to be small whenever it is less than magnitude specific distance, *r0*, which is given (in kilometres) by:

$$
r0 = 10^{0.5 \, M_w - 2.25}
$$
 Equation 2.6

While simple and quick, this method does not take into consideration factors such as bathymetry or source orientation. It is thus subject to considerable uncertainty. This is taken into consideration in an approximate way in the calculation of an 'effective magnitude' *Meff* in NTHM 2013 (see Section 6.5 of Power [2013]). *Meff* is an 'effective magnitude' which takes into account a range of different causes of uncertainty as if they had an effect equivalent to increasing or decreasing the earthquake magnitude.

For NTHM 2021, we decided to use hydrodynamic modelling to calculate the heights of tsunami generated by local crustal sources at their characteristic magnitudes and then use the method described in Section [2.2](#page-9-0) (i.e. calculating a source and coastal zone specific *Bij* and standard deviation) to calculate the tsunami shoreline wave heights for other earthquake magnitudes. To do this, we selected local crustal sources that were offshore and capable of producing an earthquake above M_W 6.5 according to the local faults used in Stirling et al. (2012). We also included any onshore local source in that study capable of generating an M_W 7.5 earthquake or greater, or M_W 8.0 if the mechanism was strike-slip. [Figure 2.3](#page-12-0) shows the location of the 250 local crustal sources used in this study.

Figure 2.3 Map showing the local crustal sources used in NTHM 2021.

For each local crustal source, we then estimated the tsunami heights by running a COMCOT model using the preferred characteristic magnitude. Details of the magnitudes and recurrence intervals are given in Appendix 1.

Each tsunami scenario was modelled using the COMCOT tsunami simulation model on a series of nested grids (Wang and Power 2011), using the same grid configuration as was used for the tsunami scenario database (Gusman et al. 2019, 2020). The nested grid elevation data and coverages are shown in [Figure 2.2.](#page-11-1) The DEM used to build the modelling grids are the same as those used in Gusman et al. (2019, 2020). The boundaries of the modelling grids are shown in [Table 2.2.](#page-13-3) Each model was run for 10 hours. Vertical wall boundaries are implemented at the

10 cm water-depth contour. The slip amount of each local source scenario was calculated from the earthquake magnitude by assuming a rigidity of 34.3 GPa. Then, the ground surface and seafloor displacement were calculated using the elastic theory documented in Okada (1985).

Grid Layer	West	East	South	North
	150,0000	200.00	-55.0000	-25.0000
	160,0000	190,0000	-50.0000	-30.0000
ິ Q	166.0000	179,0000	-48.0000	-34.0000
	182.5000	184.5000	-45.0000	-43.0000

Table 2.2 Boundaries of the nested grids.

Figure 2.4 Nested grid modelling set-up used for the NTHM 2021 update. See Gusman et al. (2019, 2020) for details.

In total, 250 tsunami scenarios were used to inform the probabilistic model, one for each of the local fault sources. For each scenario *i* (here representing a particular fault), a set of *Bij* values were calculated using the tsunami heights in each coastal region *j* using [Equation 2.2](#page-10-3) and the preferred characteristic magnitude for that fault, M_W. For the probabilistic tsunami hazard calculation, the *Bij* values were used to calculate the shoreline wave heights at each coastal zone for every earthquake in a particular synthetic catalogue.

2.5 Updates to the Tentatively Identified Local Faults

The 2013 NTHM (Power 2013) tentatively identified several sets of local faults that could be tsunamigenic. These fell into three groups: Outer Rise Faults, Taranaki Basin Faults and Offshore West Coast Faults. Since that time, research has changed our understanding of some of these faults, and here we provide details of the updated representations of these faults as used in the 2021 NTHM.

2.5.1 Outer Rise Faults

Bathymetric mapping of the seafloor east of the southern Hikurangi Trench has not identified evidence of outer rise faults. It is suspected that, due to the thickness and rigidity of the incoming Hikurangi Plateau, substantial bending of the incoming plate does not occur until after the point at which the incoming plate has started to subduct. In the northern half of the Hikurangi subduction zone, earthquakes caused by normal faulting in the subducting slab are fairly common, and we suggest these are playing the role that is played by Outer Rise faults in other subduction zones (i.e. associated with bending of the incoming plate). If these

earthquakes are large enough, they could still cause sufficient deformation to be tsunamigenic. In the southern half of the Hikurangi subduction zone, normal faulting earthquakes in the subducting plate are much less frequently observed. We suggest that this reflects differences in the coupling between plates along the margin – in the north, the coupling is weak and movement of the plates is continuously applying bending forces to the slab, whereas, in the south, the strong interseismic coupling appears to only be released in large infrequent earthquakes, and we may anticipate that the associated normal faulting in the slab tends to occur in the period following one of these large events.

Further research is highly desirable to improve our understanding of these faults and their potential to generate tsunami. For the purposes of the 2021 NTHM, we have adapted the representation of these faults as shown in [Figure 2.5](#page-15-0) and [Table 2.3.](#page-14-0) For consistency with the 2013 NTHM, we continue to label these as 'Outer Rise Faults', even though they would be more accurately called 'Normal Faults in the Subducting Slab'.

Name	Type	Length (km)	Dip	Dip Direction	Depth of Top (km)	Depth of Bottom (km)	M_{W}	Recurrence Interval (Years)
Raukumara Outer Rise	nn	150	58°	300°	5	30	8.0	1300
Hawkes Bay Outer Rise	nn	150	58°	300°	5	30	8.0	1460
North Wairarapa Outer Rise	nn	150	58°	308°	5	30	8.0	1640
South Wairarapa Outer Rise	nn	150	58°	315°	5	30	8.0	1900

Table 2.3 Assumed Hikurangi Outer Rise fault properties. Type 'nn' implies a normal fault mechanism.

Figure 2.5 Assumed location of Hikurangi 'Outer Rise' faults as used for this study.

2.5.2 Taranaki Basin Faults

Recent research by Seebeck et al. (2021) has improved our understanding of these faults. The revised locations are shown in [Figure 2.6,](#page-16-1) and properties are shown in [Table 2.4.](#page-16-2) What was formerly called the 'Manaia South Fault' is now called the 'Waimea Offshore Fault' for consistency with the naming used by the Community Fault Model (CFM; Rattenbury 2020). The 'Manaia North Fault' in the 2013 NTHM has been removed, as it is regarded as inactive.

Figure 2.6 Assumed locations of tentatively identified Taranaki Basin faults.

Table 2.4 Assumed Taranaki Basin fault properties. Type 'nn' implies a normal fault mechanism, 'rv' a reverse mechanism, 'rs' a combined reverse and strike-slip.

2.5.3 Offshore West Coast Faults

The representation of faults offshore of the west coast of the South Island has been revised, largely toward consistency with the CFM. Note in particular the re-numbering of the Cape Foulwind fault segments. The revised locations are shown in Figures [2.7](#page-17-1) and [2.8,](#page-18-0) and properties are shown in [Table 2.5.](#page-19-1)

Figure 2.7 Assumed locations of west coast South Island faults (northern view).

Figure 2.8 Assumed locations of west coast South Island faults (southern view).

Name	Type	Length (km)	Dip	Dip Direction	Depth of Top (km)	Depth of Bottom (km)	Mw	Recurrence Interval (Years)
Barn	rv	67	30°	143°	0	15	7.6	2400
SouthWestland1	rv	65	45°	126°	0	15	7.6	9700
SouthWestland2	rv	65	45°	143°	0	15	7.6	4900
SouthWestland3	rv	65	45°	140°	$\mathbf 0$	15	7.6	4900
SouthWestland4	rv	65	45°	150°	0	15	7.6	4900
CapeFoulwind4	rv	65	60°	139°	$\mathbf 0$	15	7.6	9700
CapeFoulwind1	rv	86	60°	105°	0	15	7.6	9700
Kongahu	rv	96	60°	90°	$\mathbf 0$	15	7.6	15,000
Kahurangi	rv	103	50°	118°	0	15	7.6	15,000

Table 2.5 Assumed west coast South Island fault properties. Type 'rv' implies a reverse fault mechanism.

2.6 Other Modifications from National Tsunami Hazard Model 2013

Throughout the tsunami scenario modelling of subduction zone sources, and the probabilistic calculations of the hazard posed by them, a rigidity of 40 Gpa was assumed (in contrast to 50 GPa in NTHM 2013). For the scenario modelling of crustal faults, a rigidity of 34.3 GPa was assumed (Burbidge et al. 2021).

To improve the statistical quality of the estimation of uncertainty, the number of samples of epistemic uncertainty (see Section 6 of Power [2013]) was increased from 300 to 600.

In the deaggregation calculation (see Section 6.8 of Power [2013]), instead of selecting the three closest events from each synthetic catalogue, the 10 closest were selected, this was done to improve the statistical robustness of the deaggregation.

3.0 RESULTS

3.1 National Tsunami Hazard Maps 2021

Figures [3.1](#page-21-0)[–3.4](#page-24-0) show the shoreline^{[3](#page-20-2)} tsunami hazard in NTHM 2021 with the improvements described in Section [2.](#page-7-0) Maps show the shoreline hazard at four return periods, 100, 500, 1000 and 2500 years, for the tsunami hazard in each coastal zone. Here, the tsunami hazard level is shown at the $50th$ (best estimate) and $84th$ (conservative estimate) percentile of epistemic uncertainty in each coastal section.

Note that, in the following figures, we refer to the 'maximum' tsunami amplitude in each zone. Strictly speaking, this is the 99th percentile of the tsunami maximum amplitude for all coastal points in each hazard zone polygon; this reduces the effect of very localised outliers.

³ The numerical tsunami modelling that underpins the assessment of shoreline tsunami heights has been created using innermost nested modelling grids of 15 arc-second resolution (310–380 m) and assuming vertical-wall boundary conditions at the 10 cm depth contour. 'Shoreline' tsunami heights are approximated by the maximum water levels in the grid cells adjacent to the wall boundary.

Tsunami Height (Maximum Amplitude) in metres at 50th percentile at return period:100

Tsunami Height (Maximum Amplitude) in metres at 84th percentile at return period:100

Figure 3.1 Expected maximum tsunami height in metres at the 100-year return period, shown at median (50th) and 84th percentiles of epistemic uncertainty.

Tsunami Height (Maximum Amplitude) in metres at 50th percentile at return period:500

Tsunami Height (Maximum Amplitude) in metres at 84th percentile at return period:500

Figure 3.2 Expected maximum tsunami height in metres at the 500-year return period, shown at median (50th) and 84th percentiles of epistemic uncertainty.

Tsunami Height (Maximum Amplitude) in metres at 50th percentile at return period:1000

Tsunami Height (Maximum Amplitude) in metres at 84th percentile at return period:1000

Figure 3.3 Expected maximum tsunami height in metres at the 1000-year return period, shown at median (50th) and 84th percentiles of epistemic uncertainty.

Tsunami Height (Maximum Amplitude) in metres at 50th percentile at return period:2500

Tsunami Height (Maximum Amplitude) in metres at 84th percentile at return period:2500

Maps showing the difference between NTHM 2021 and NTHM 2013 at the 500-year return period (50% level of confidence) are shown in [Figure 3.5.](#page-25-1) At this return period, the estimated tsunami hazard is mostly lower in the 2021 study, which we interpret as being mostly due to the more accurate modelling of distant tsunami sources in the revised model.

Figure 3.5 Comparison of the 500-year tsunami hazard (at the 50% level of confidence) in the 2013 (left) and 2021 (right) tsunami hazard models.

At the 2500-year return period (50% level of confidence) shown in [Figure 3.6,](#page-25-2) we see a reduction in the estimated hazard along the west coast of the South Island and in the Bay of Plenty but an increase in the estimated hazard along the east coast of the North Island. The increased hazard along the east coast of the North Island we interpret as being related to the more accurate modelling applied to the 'Outer Rise' sources (see Section 2.5.1).

Figure 3.6 Comparison of the 2500-year tsunami hazard (at the 50% level of confidence) in the 2013 (left) and 2021 (right) tsunami hazard models.

3.2 Tsunami Hazard Curves and Hazard Deaggregation for Major Cities

In this section, we show the outline of the coastal section, the tsunami hazard curves for return periods between 100 and 2500 years, the deaggregation of the tsunami hazard at the 500 year return period and the deaggregation of the hazard at the 2500-year return period for each of the major coastal cities. The fault name codes used in the deaggregation plots are given in Appendix 1.

Auckland East Coast

Figure 3.7 Area map and tsunami hazard curve for Auckland East.

Deaggregation of Zone:31, Return Period:500 years, Tsunami Height (Maximum Amplitude) at:50th percentile:2.405 m

Deaggregation of Zone:31, Return Period:2500 years, Tsunami Height (Maximum Amplitude) at:50th percentile:4.006 m

Figure 3.8 Deaggregation of tsunami sources for Auckland East Coast at the 500-year (top) and 2500-year (bottom) return periods.

Auckland West Coast

Figure 3.9 Area map and tsunami hazard curve for Auckland West.

Deaggregation of Zone:124, Return Period:500 years, Tsunami Height (Maximum Amplitude) at:50th percentile:1.765 m

Deaggregation of Zone:124, Return Period:2500 years, Tsunami Height (Maximum Amplitude) at:50th percentile:2.87 m

Figure 3.10 Deaggregation of tsunami sources for Auckland West Coast at the 500-year (top) and 2500-year (bottom) return periods.

Christchurch

Figure 3.11 Area map and tsunami hazard curve for Christchurch.

Deaggregation of Zone:150, Return Period:500 years, Tsunami Height (Maximum Amplitude) at:50th percentile:4.377 m

Deaggregation of Zone:150, Return Period:2500 years, Tsunami Height (Maximum Amplitude) at:50th percentile:6.501 m

Figure 3.12 Deaggregation of tsunami sources for Christchurch at the 500-year (top) and 2500-year (bottom) return periods.

Dunedin

Figure 3.13 Area map and tsunami hazard curve for Dunedin.

Deaggregation of Zone:170, Return Period:500 years, Tsunami Height (Maximum Amplitude) at:50th percentile:2.82 m

Deaggregation of Zone:170, Return Period:2500 years, Tsunami Height (Maximum Amplitude) at:50th percentile:4.453 m

Figure 3.14 Deaggregation of tsunami sources for Dunedin at the 500-year (top) and 2500-year (bottom) return periods.

Gisborne

Figure 3.15 Area map and tsunami hazard curve for Gisborne.

Deaggregation of Zone:65, Return Period:500 years, Tsunami Height (Maximum Amplitude) at:50th percentile:6.155 m

Deaggregation of Zone:65, Return Period:2500 years, Tsunami Height (Maximum Amplitude) at:50th percentile:10.76 m

Figure 3.16 Deaggregation of tsunami sources for Gisborne at the 500-year (top) and 2500-year (bottom) return periods.

Invercargill

Figure 3.17 Area map and tsunami hazard curve for Invercargill.

Deaggregation of Zone:181, Return Period:500 years, Tsunami Height (Maximum Amplitude) at:50th percentile:4.049 m

Deaggregation of Zone:181, Return Period:2500 years, Tsunami Height (Maximum Amplitude) at:50th percentile:6.876 m

Figure 3.18 Deaggregation of tsunami sources for Invercargill at the 500-year (top) and 2500-year (bottom) return periods.

Kāpiti Coast

Figure 3.19 Area map and tsunami hazard curve for Kāpiti Coast.

Deaggregation of Zone:95, Return Period:500 years, Tsunami Height (Maximum Amplitude) at:50th percentile:2.433 m

Deaggregation of Zone:95, Return Period:2500 years, Tsunami Height (Maximum Amplitude) at:50th percentile:4.462 m

Figure 3.20 Deaggregation of tsunami sources for Kāpiti Coast at the 500-year (top) and 2500-year (bottom) return periods.

Napier

Figure 3.21 Area map and tsunami hazard curve for Napier.

Deaggregation of Zone:73, Return Period:500 years, Tsunami Height (Maximum Amplitude) at:50th percentile:6.199 m

HawkesBayOuterRise_1002:8.13_231:7.8

Deaggregation of Zone:73, Return Period:2500 years, Tsunami Height (Maximum Amplitude) at:50th percentile:11.28 m

Figure 3.22 Deaggregation of tsunami sources for Napier at the 500-year (top) and 2500-year (bottom) return periods.

Nelson

Figure 3.23 Area map and tsunami hazard curve for Nelson.

Deaggregation of Zone:247, Return Period:500 years, Tsunami Height (Maximum Amplitude) at:50th percentile:2.487 m

Deaggregation of Zone:247, Return Period:2500 years, Tsunami Height (Maximum Amplitude) at:50th percentile:3.987 m

Figure 3.24 Deaggregation of tsunami sources for Nelson at the 500-year (top) and 2500-year (bottom) return periods.

New Plymouth

Figure 3.25 Area map and tsunami hazard curve for New Plymouth.

Deaggregation of Zone:111, Return Period:2500 years, Tsunami Height (Maximum Amplitude) at:50th percentile:4.1 m

Figure 3.26 Deaggregation of tsunami sources for New Plymouth at the 500-year (top) and 2500-year (bottom) return periods.

Porirua

Figure 3.27 Area map and tsunami hazard curve for Porirua.

Deaggregation of Zone:94, Return Period:2500 years, Tsunami Height (Maximum Amplitude) at:50th percentile:5.08 m

Figure 3.28 Deaggregation of tsunami sources for Porirua at the 500-year (top) and 2500-year (bottom) return periods.

Tauranga

Figure 3.29 Area map and tsunami hazard curve for Tauranga.

Deaggregation of Zone:47, Return Period:500 years, Tsunami Height (Maximum Amplitude) at:50th percentile:3.335 m

Deaggregation of Zone:47, Return Period:2500 years, Tsunami Height (Maximum Amplitude) at:50th percentile:5.478 m

Figure 3.30 Deaggregation of tsunami sources for Tauranga at the 500-year (top) and 2500-year (bottom) return periods.

Timaru

Figure 3.31 Area map and tsunami hazard curve for Timaru.

Deaggregation of Zone:161, Return Period:500 years, Tsunami Height (Maximum Amplitude) at:50th percentile:2.836 m

Deaggregation of Zone:161, Return Period:2500 years, Tsunami Height (Maximum Amplitude) at:50th percentile:4.256 m

Figure 3.32 Deaggregation of tsunami sources for Timaru at the 500-year (top) and 2500-year (bottom) return periods.

Wellington

Figure 3.33 Area map and tsunami hazard curve for Wellington.

Deaggregation of Zone:91, Return Period:500 years, Tsunami Height (Maximum Amplitude) at:50th percentile:4.61 m

Deaggregation of Zone:91, Return Period:2500 years, Tsunami Height (Maximum Amplitude) at:50th percentile:8.205 m

Figure 3.34 Deaggregation of tsunami sources for Wellington at the 500-year (top) and 2500-year (bottom) return periods.

Whakatāne

Figure 3.35 Area map and tsunami hazard curve for Whakatāne.

Deaggregation of Zone:50, Return Period:2500 years, Tsunami Height (Maximum Amplitude) at:50th percentile:5.695 m

Figure 3.36 Deaggregation of tsunami sources for Whakatāne at the 500-year (top) and 2500-year (bottom) return periods.

Whangārei

Figure 3.37 Area map and tsunami hazard curve for Whangārei.

Deaggregation of Zone:20, Return Period:500 years, Tsunami Height (Maximum Amplitude) at:50th percentile:3.535 m

Figure 3.38 Deaggregation of tsunami sources for Whangārei at the 500-year (top) and 2500-year (bottom) return periods.

4.0 DISCUSSION AND LIMITATIONS

The 2021 updates to the NTHM should significantly improve the accuracy of the hazard estimates, primarily because of the improved accuracy with which the tsunami heights have been estimated for the events in our synthetic catalogue of earthquakes.

For the most part, the hazard estimates appear to have been revised lower, although they remain broadly consistent with our knowledge of historical and paleo-tsunami (for instance, when compared with Section 3 of Power [2013]). Where hazard curves are now lower, the modelling improvements will also lead to better selection of scenarios for deaggregation, with the net effect that estimates of inundation extents (where these have previously been made using deaggregations of the 2013 NTHM) will not necessarily decrease, as the scenarios selected by the improved deaggregation may still be able to inundate further inland when modelled from source through to inundation.

The 2021 model does show increased estimates of hazard for the east coast of the North Island at long return periods. We interpret this as being due to better estimation of tsunami heights caused by earthquakes in the subducting slab (these are also referred to as 'outer rise' earthquakes for consistency with the naming used in the 2013 study). As these faults are not at all well characterised, and not even proven to exist along the southern part of the Hikurangi margin, we see this as an important area of uncertainty in our results and would like to encourage study of the potential for large earthquakes in the slab as a research topic for the coming years.

As with the 2013 study, the 2021 probabilistic tsunami hazard model represents the best endeavours of the report authors at the time it was created. Scientific understanding of input parameters will continue to evolve, and improved methods for calculating the hazard will be developed. The programs used to perform the calculations are complicated, and programming errors may be found and corrected. Hence, the results in this report represent only a snapshot of the estimated tsunami hazard as determined at the time of its construction.

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APPENDICES

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APPENDIX 1 CRUSTAL FAULT PARAMETERS, AS USED IN THE 2021 NATIONAL TSUNAMI HAZARD MODEL

Table A1.1 Crustal fault properties. Fault Name and NZSHM_Number are as used in the National Seismic Hazard Model (Stirling et al. 2012). NZSHM_Number can be used to identify fault locations using the figures in Stirling et al. (2012). MWMN, MW, MWMX are minimum, preferred and maximum moment magnitudes. RECINTMN, RECINT and RECINTMX are minimum, preferred and maximum recurrence intervals (in years).

⁴ The Wharekauhau Fault is assumed in the 2021 National Tsunami Hazard Model to not rupture independently of the Wairarapa Fault; instead, 50% of Wairarapa Fault ruptures are assumed to also include rupture of the Wharekauhau Fault.

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